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WATER RESOURCES OF DELAWARE COUNTY, INDIANA

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WATER RESOURCES OF DELAWARE COUNTY, INDIANA

By R. E. Hoggatt, J. D. Hunn,

and W. J. Steen

ABSTRACT

In Delaware County, Indiana, water of sufficient quantity and quality is available with proper development in the streams and in the aquifers for the expected increase of most agricultural, domestic, industrial, and municipal uses in the foreseeable future.

Water use in Delaware County totals about 18.5 mgd (million gallons per day) of which 63 percent is from surface water and 37 percent from ground water. Muncie uses about 80 percent of the total water used in the county.

The White River at Muncie yields a natural flow (after adjustment for diversion) of more than 20 cfs (cubic feet per second) (13 mgd) 90 percent of the time from a drainage area of 242 square miles. Low flow in White River is augmented by releases from Prairie Creek Reservoir for the Muncie water supply.

The Mississinewa River offers good development potential, although the low flow is not as well sustained as in the White River. By way of comparison, the Mississinewa River 3 miles upstream from Eaton, drainage area 304 square miles, yields more than 8 cfs (5 mgd) 90 percent of the time.

Many small streams in the northern half of the county frequently go dry, while most streams in the southern half of the county, regardless of size, yield more than 0.013 mgd per square mile 90 percent of the time. Buck Creek has the highest low-flow yield of all the small streams in the county. At the stream-gaging station on Buck Creek near Muncie, drainage area 36.7 square miles, the flow exceeds 10 cfs (6.5 mgd) 90 percent of the time.

Nearly all the available ground water occurs within 400 feet of the land surface. Wells utilizing all available water-bearing zones could probably yield as much as 1,000 gpm (gallons per minute) in the south central and southeastern parts of the county.

The upper 100 feet of dolomite of Silurian Age, at the bedrock surface, is one of the two principal aquifers in the county. It underlies the entire county at depths ranging from 0 to about 300 feet. The most permeable parts of this aquifer could yield as much as 500 gpm to individual wells.

Sand and gravel in and near buried valleys yield water for domestic and industrial use. The areal extent and potential of this unit are not well known.

The principal Pleistocene aquifer is composed chiefly of glacio-fluvial sand and gravel up to 80 feet thick and underlies most of the county at depths of 0 to about 140 feet. The thicker parts of this aquifer could yield as much as 500 gpm to individual wells.

The surface and ground waters in Delaware County are generally a calcium bicarbonate type with moderate amounts of sulfate and hardness. The ground water may contain objectionable concentrations of iron. The chemical quality of the ground water is similar in all important aquifers in the county. The quality of the streams at about 90 percent flow duration approximates that of the ground water except where activities of man have increased the mineralization of the streams.

PURPOSE AND SCOPE

The purpose of this report is to describe for those concerned with water development and use in Delaware County (1) the sources of available water for current and future uses, (2) the quantity, distribution, and quality of the resource, and (3) some of the significant relations in the hydrologic system between precipitation, geology, ground water, surface water, and water quality.

This purpose was accomplished by analyzing information based on (1) well data and aquifer characteristics collected from drillers, water works superintendents, industries and others, (2) streamflow data based on 7 stream-gaging station records, 4 partial-record station measurements and base-flow measurements at miscellaneous sites, (3) quality-of-water data based on samples collected from the streams during low flows, from wells, and at water quality monitoring stations, and (4) field investigation of the surficial geology of the county.

This report provides information that will be useful in planning the development, control, and utilization of surface and ground waters in Delaware County. Some of the problems to be considered by the water user in the development of the resource are discussed.

COOPERATION AND ACKNOWLEDGEMENTS

This report was prepared as part of a cooperative program between the Indiana Department of Natural Resources, Division of Water and the U. S. Geological Survey, Water Resources Division. The work was done under the general guidance of the Indiana Council, Water Resources Division, U. S. Geological Survey.

The authors thank all persons and State Agencies who contributed time, information, and assistance during the preparation of this report. We especially thank the well drillers who furnished well logs and other

information, as well as the Muncie Water Works Company, industries, firms, and water works superintendents who contributed information. The Indiana State Highway Commission supplied locations and logs of test borings and the Indiana State Board of Health and Stream Pollution Control Board furnished water-quality data at monitoring stations on the Mississinewa and White Rivers.

WATER USE

Water use in Delaware County is presently about 18.5 mgd (million gallons per day) or 155 gpd (gallons per day) per person. Water use has been increasing at the rate of about 2 gallons per person per year since 1950. It is estimated that by 1980 the requirements will be nearly 190 gpd per person or a total of 31.5 mgd for the county.

An inventory of municipal and industrial water consumption was made in order to obtain information on present water use in the county and to outline future water needs. These data and records, obtained from water users, the Indiana Public Service Commission, Indiana State Board of Health, and other sources, were utilized in formulating an estimate of the amount of water used in the county.

A compilation of these data along with projected estimates of water use in 1980 are shown in table 1. The estimates for 1980 are made on the assumption that the use of surface water and ground water will have the same ratio as at the present time and that Muncie will be the only surface water user. A small amount of surface water has been used in time past for irrigation, probably less than 0.1 mgd. A larger amount will probably be used in the future when climatic factors are unfavorable during the growing season. As water used for irrigation is such a nebulous figure to estimate in Delaware County, no estimate has been entered in table 1 for either 1964 or 1980.

Table 1. Estimates of water use in Delaware County in millions of gallons per day, for 1964 and 1980.

Source and use	1964	1980
Surface Water		
Industrial	5.8	9.8
Municipal	5.9	10.0
Self-supplied domestic	0	0
Total	11.7	19.8
Ground Water		···-
Industrial	3.7	6.4
Municipal	.6	1.0
Self-supplied domestic	2.5	4.3
Total	6.8	11.7
Grand Total	18.5	31.5

Water use today is the highest in the history of Delaware County, and the trend has been upward for many years with no indication of leveling off. Of the 18.5 mgd used at present, industry uses about one-half, with the remaining part divided between municipal (Albany, Eaton, Gaston, Muncie and Yorktown) and self-supplied domestic use (table 1). Most municipal and industrial use is not consumptive, and used water is disposed of in the streams usually after treatment.

Industrial and municipal use is concentrated in the Muncie area, where the combined use is about 14.7 mgd, or 80 percent of the total water used in the county. This 14.7 mgd includes water taken from White River, Prairie Creek Reservoir, Buck Creek, and wells. Of the total used in the county, about 11.7 mgd, or 63 percent, is taken from surface water sources, with the remainder, 37 percent, being obtained from numerous municipal, industrial and domestic ground-water supplies.

Population trends have likewise been progressing upward. With an expected increase in population from 120,000 (estimated) in 1964 to about 167,000 by 1980 (Hannaford, 1960, p. 7), water consumption is expected to increase from the present 155 gpd per person (18.5 mgd) to about 190 gpd per person (31.5 mgd). Of this estimated total volume, about 19.8 mgd will probably be taken from surface water sources, primarily White River, Prairie Creek Reservoir and Buck Creek, and the remainder from well supplies.

PHYSICAL FEATURES AFFECTING WATER SUPPLY

Climate

Precipitation, temperature, and wind are climatic factors directly affecting surface and ground-water supplies. Precipitation adds water to the streams and ground-water reservoirs causing water levels to rise. During dry periods water used or flowing out of an area is not replenished, and water levels fall. High temperature and wind cause water levels to fall by increasing the amount of water returned to the atmosphere by evaporation (from exposed water surfaces) and transpiration (from plants). High temperatures likewise cause water levels to fall by increasing the withdrawals by man for both his personal comfort and industry.

In east-central Indiana, which embraces Delaware County, precipitation averages about 39 inches per year, with monthly averages ranging from 2.37 inches for February to 4.50 inches for June. Much of the precipitation that replenishes the ground-water supply occurs during the early spring and late fall months. Although precipitation is slightly less during these periods, recharge is greater than during the growing season, when monthly precipitation averages are higher. During the growing season, the evaporation and the transpiration losses are large and percolation of precipitation to the ground-water reservoirs is greatly decreased. The average monthly, annual, and seasonal (April through September) precipitation and temperature of east-central Indiana for the period 1931-60 are shown in table 2.

Table 2. Average monthly, annual, and seasonal (April through September) precipitation and temperature of east-central Indiana, 1931-60.

Month	Average Precipitation (inches)	Average Temperature (^O F)
January	3.03	29,4
February	2.37	30.9
March	3.35	38.8
April	3.78	50.5
May	4.13	60.9
June	4.50	70.7
July	3.60	74.2
August	2,98	72.5
September	3,31	65.5
October	2.77	54.5
November	2,93	41.1
December	2.52	31.1
Annuel	39.27	51.7
April through September	22,30	65.7

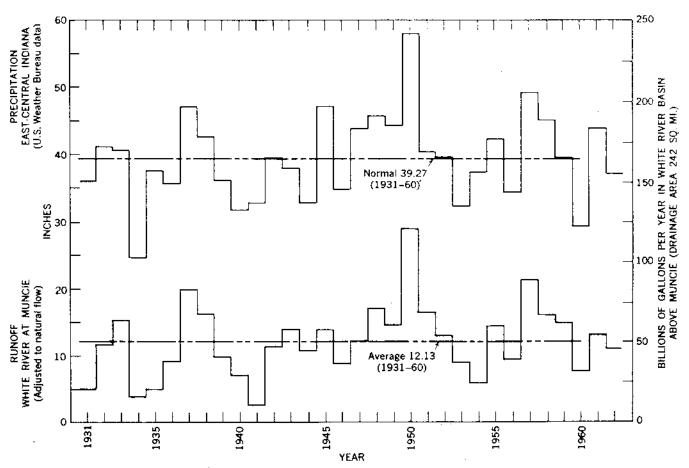
From U. S. Weather Bureau Data

The annual runoff of the White River at Muncie, adjusted to natural flow, is compared with the annual precipitation of east-central Indiana for the period of 1931-62 in figure 1. The similarity between the two graphs (A of fig. 1) is apparent, with the lowest annual flows in White River generally coinciding with periods of lowest precipitation. The ratio of runoff to precipitation is not constant, however, (B of fig. 1) varying from 0.08 in 1941 to 0.50 in 1950. Soil moisture, amount of precipitation, intensity of precipitation, distribution of precipitation throughout the year, and water levels in ground-water aquifers, as well as other factors, have their effect on the proportion of precipitation that runs off into the stream by either overland or underground flow.

Topography and Drainage

The topography of Delaware County is flat to gently rolling with a dendritic drainage pattern.

Land elevations are highest in the southeastern part of the county where the maximum is about 1,100 feet above mean sea level. The land slopes downward gradually from this area towards the northwest and west. The lowest elevation is 835 feet above mean sea level in the northwestern part where the Mississinewa River flows from the county. Maximum local relief of approximately 80 feet occurs near the south-central edge of the county.



A. The lowest annual flow in White River generally coincides with the lowest annual precipitation

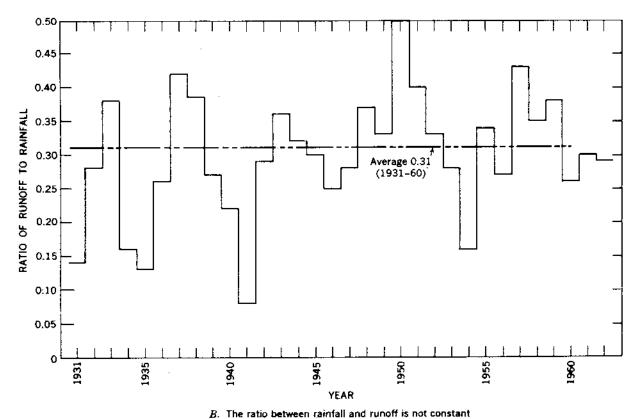


Figure 1.—About one-third of the precipitation that falls on Delaware County

runs off in the streams.

Delaware County lies within the White River and Mississinewa River drainage basins. Of the 392 square miles of area in the county, approximately two-thirds is drained by the White River system which includes Prairie, Buck, York Prairie, Killbuck and Pipe Creeks. The Mississinewa River and its tributaries drain the northeastern part of the county.

The White River flows generally westward through the county, and the Mississinewa flows in a northwesterly direction. The slopes of the stream channels are about 3.3 and 3.9 feet per mile for the Mississinewa and White Rivers, respectively, as they flow through the county. Several of the smaller streams have been dredged to improve the surface drainage.

Surficial Geology

The land surface of Delaware County is almost entirely covered by a relatively impermeable layer of glacial till and some silt. This deposit limits the amount of ground-water recharge and the amount of ground-water discharge to streams. The unit helps prevent the direct contamination of aquifers by man-made wastes.

Differences in base flow of the streams of Delaware County are caused by: (1) the unequal distribution and hydraulic properties of sand and gravel deposits beneath the surface till, (2) variations in thickness and permeability of the till, and (3) variations of the heights of the piezometric surfaces of the till and of the underlying saturated material above or below the level of the streams. The pick up of base flow in the streams compared with the size of the drainage area indicates the relative availability of ground water in the drift upstream from the points of measurement.

SOURCES OF WATER SUPPLY

The major sources of water in Delaware County are the streams, the upper 400 feet of the glacial drift, and the bedrock. The streams supply the largest amount of water and this amount could be increased by impoundments. The rocks generally provide adequate amounts of water and this amount could be increased by additional wells.

Streams

The average annual precipitation for the county is 39 inches. Of this amount about 12 inches, or 82 billion gallons, drains out of the county through the stream system. The quantity of water in the streams at any given time varies considerably from flood to drought. The amount that can be obtained as a dependable supply is highly variable from stream to stream, as well as at different locations on the same stream.

Application of Streamflow Data

The water available in the streams without impoundment has been evaluated by three techniques: (1) flow duration, (2) low-flow frequency, and (3) base-flow discharge measurements. The water that would be available by impoundment in the vicinity of the stream-gaging stations has been evaluated by a fourth technique: determination of draft-storage requirements. The first two techniques are statistical studies based on records of daily discharges collected at stream-gaging stations. The third technique is the comparison of base-flow (ground-water runoff) increase or decrease from various streams by obtaining discharge measurements when there is no overland runoff into the streams. The fourth technique is based on the low-flow frequency analysis and is used as a design tool. The two statistical studies were developed for a common time interval of 1924-63 by correlation of short-term with long-term water discharge records. A common reference period tends to equalize or make homogeneous climatic experiences.

These studies were based on records from 7 stream-gaging stations and measurements at 4 partial-record stations and 69 miscellaneous sites (including some observations of no flow) in the upper Mississinewa and White River basins embracing Delaware County. A list of the stations and sites is given in table 3, and the location of each is shown on plate 1.

The works of man have greatly altered the discharge of some streams or certain reaches of the channels. Most of these man-made effects are much more pronounced during low flows. At these times a considerable percentage of the total flow may be man supplied through waste or sanitary systems and reservoir releases, or man deleted or detained, through water supplies including reservoir storage. Adjustments have been made to some of the observed data to approach natural flow. Where observed data have been adjusted, this has been noted on the figures and in the tables.

Flow duration. Flow characteristics of a stream throughout its range in discharge may be shown by the flow-duration curve. A flow-duration curve shows the percentage of time discharges have been equaled or exceeded during a given period without regard to the sequence of occurrence. If the period upon which the curve is based represents a long-term flow of a stream, the curve may be used to estimate distribution of future flows for water supply and pollution studies. The slope of the duration curve is a quantitative measure of the variability of the streamflow and a flat slope on the lower end depicts well sustained flow.

The variation in flow at the 7 gaging stations is illustrated in figure 2. These flow-duration curves have been expressed in cubic feet per second per square mile (cfsm) so that they might be compared on a unit basis.

Table 3,--List of stream-gaging stations, partial-record stations, and miscellaneous sites used in this report and shown on Plate 1

K. S.	5,42	Brandon ditch near Middleton	. s. 80	6.27 M.	Mud Creek near Selma	40
H. S.	4.97	Fall Creek near Cross Roads	S.	179 M.	White River near Selma	39
s. G.	814	White River near Noblesville	. s. 78	1.03 H	Stoney Creek tributary near Blountsville	38
M. S.	108	Pipe Creek near Frankton	. s. 77	18.7 M.	Stoney Creek near Blountsville	37
P. R.	44.7	Pipe Creek near Alexandria	. s. 76	84.5 H.	White River near Farmland	36
H. S.	18.3	Pipe Creek near Gaston	. S. 75	8.06 M	Eightmile Creek near Maxville	35
H. S.	.67	Steel ditch near Gaston	. S. 74	1.15	White River tributary near Maxville	34
M. S.	7.43	Yeager, Finley, Manard ditch near Gaston	l. S. 73	34.1 M.	White River at Winchester	33
M. S.	1.30	Pipe Creek at Gaston	P. R. S. 72	21.3 F	White River near Harrisville	32
s. S.	557	White River at Perkinsville	i. G. S. 71	677 S	Mississinewa River at Marion	31
P. R.	98.0	Killbuck Creek near Anderson	M. S. 70	515	Mississinewa River near Upland	30
¥. S.	74.4	Killbuck Creek near North Anderson	M. S. 69	20.7	Barren Creek near Upland	29
K. S.	4.91	Pleasant Run Creek near Gilman	M. S. 68	9.15	Lake Branch near Upland	28
×. S.	18,5	Jakes Creek near Bethel	H. S. 67	8.14	Hoppas ditch at Matthews	27
M. S.	5.01	Eagle Branch near Cammack	M. S. 66	21.5	Pike Creek at Wheeling	26
×.	8,64	Jakes Creek near Cammack	M. S. 65	2.76	Hedgeland ditch near Wheeling	25
M. S.	1,63	Thurston ditch near Bethel	н. s. 64	4.45	Hayden ditch near Wheeling	24
M. S.	26.1	Killbuck Creek near Bethel	M. S. 63	10.5	Studebaker ditch near Wheeling	23
×.	6.52	Mud Creek near Anthony	P. R. S. 62	76.4	Bick Lick Creek near Wheeling	22
M. S.	7.35	Killbuck Creek near Anthony	M. S. 61	32.9	Big Lick Creek near Hartford City	21
S. G.	401	White River at Anderson	H. S. 60	348	Mississinewa River near Wheeling	20
M. S.	383	White River near Daleville	K. S. 59	13.5	Rees ditch near Eaton	19
K.S.	16.2	York Prairie Creek near Yorktown	¥. S. 58	8.07	Rees ditch near Millgrove	18
E S	9.70	York Prairie Creek near Cammack	M. S. 57	6.43	Rees ditch near Dunkirk	17
M. s.	3.43	York Prairie Creek near Muncie	s. G. s. 56	304	Mississinewa River near Eaton	16
M. S.	2.63	Little No Name Creek near Progress	M. S. 55	6,10	Bosman ditch near Eaton	15
M. S.	7.55	No Name Creek near Progress	H. S. 54	1.56	Bosman ditch tributary near Albany	14
ĸ.s.	11.7	Williams Creek near Cross Roads	M. S. 53	3.12	Bosman ditch near Albany	13
M. S.	16.0	Bell Creek near Cross Roads	M. S. 52	.93	Mississinewa River tributary near Royerton	12
ĸ.s.	48.5	Buck Creek near Yorktown	M. S. 51	2.35	Mississinewa River tributary near Albany	11
S. G.	36.7	Buck Creek near Muncle	M. S. 50	20.4	Campbells Creek near Desoto	10
¥. S.	27.1	Buck Creek near Oakville	M. S. 49	10.2	Campbells Creek near Parker City	ø
ĸ.s.	246	White River near Yorktown	K. S. 48	21.8	Halfway Creek near Albany	&
S. G.	242	White River at Muncle	M. S. 47	11.7	Mud Creek near Albany	7
M. S	5.50	Muncie Creek near Muncie	H. S. 46	228	Mississinewa River near Albany	6.
¥. s.	3.82	Muncie Creek near Desoto	M. S. 45	5.81	Platt Nibarger ditch near Redkey	U 1
₩. S.	3.83	White River tributary near Muncle	M. S. 44	15.0	Bear Creek near Ridgeville	4.
H. S	2.14	Medford drain near Muncie	s. G. S. 43	130	Mississinewa River near Ridgeville	ω
M. S	17.0	Prairie Creek near Muncie (Reservoir release)	M. S. 42	13.2	Harshman Creek near Salem	10
S 'W	1.95	Cunningham ditch near New Burlington	M. S. 41	25.0	Wississinewa River near Salem	juri
collection point	area (sq mi)	Stream and location	collection Reference point no.	area (sq mi)	Stream and location	Reference no.
Type of	Drainage			_		

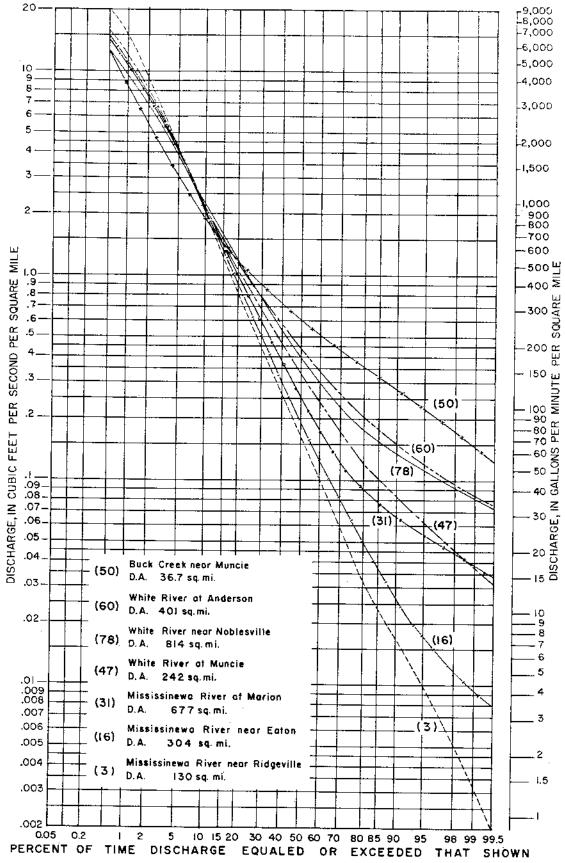


Figure 2-Duration curves of daily flows at seven gaging stations on the Mississinewa River, White River and Buck Creek (adjusted to period, 1924—1963)

Discharges of equal percent duration for the lower half of the flow-duration curves, which contain larger percentages of ground-water runoff as the flow decreases, have been tabulated and shown both in terms of cubic feet per second (cfs) and cfsm in table 4. As a means of comparing the inflow between gaging stations, table 4 also contains computed runoff figures for the intervening drainage areas.

To supplement the flow-duration data for the stream-gaging stations, the 90 and 95 percent flow-duration discharges at 4 partial-record stations and 11 miscellaneous sites have been estimated by correlation and are tabulated in table 5.

Low-flow frequency. Since the flow-duration curve does not show the frequency with which any given discharge can be expected to occur on the average, it is necessary to know more about low-flow characteristics of the stream. The low-flow frequency curve is an excellent tool for estimating how frequently a certain average low-flow discharge for a given period of consecutive days will occur. It can also be used in the analysis of storage requirements. The shapes of low-flow frequency curves, like the flow duration curves, reflect the hydrogeologic and hydraulic characteristics of the upstream basin.

Low-flow frequency data for the 3-, 7-, 14-, 30-, 60-,120-, 183-, and 274-day consecutive periods are presented in table 6 for the same 7 stream-gaging stations for which flow-duration data are tabulated and illustrated. The 7-day low-flow frequency curve for each of these stream-gaging stations is illustrated in figure 3. These low-flow frequency curves have been expressed in cfsm so that they might be compared on a unit basis.

To avoid misuse of these data it might be well to discuss recurrence interval (frequency), as used in this study. By definition, the recurrence interval, given in years, is the long-term average interval between annual minimum discharges equal to or less than the given minimum. For example, inference should not be made that the 3-day 10-year minimum will occur once, and only once during a 10-year period. It may occur in any year or successive years, but its chance of occurring as an annual 3-day minimum in any year is 1 in 10, or 10 percent.

To supplement the low-flow frequency data for the stream-gaging stations, the 7-day 2-year and 7-day 10-year low-flow have been estimated for the 4 partial-record stations and 11 miscellaneous sites where several base-flow measurements have been made and tabulated in table 5.

As with flow-duration studies, estimates of average low-flow frequency are based on records of past events and it is impossible to foresee what the future works of man will do to alter the natural flow.

Table 4.--Discharge of equal percent flow duration as taken from stream-gaging station flow-duration curves and computed for intervening areas (adjusted to the period 1924-63)

								į			
Reference	Gaging station and intervening areas	orainage area (sq mi)		Discharge per seco	of nd,	indicated per	feet p	time duration per second per	tion (cubic per square	oic feet tre mile)	<u>}</u>
			99.5	66	96	95	06	80	20	9	50
es	Mississinewa River near Ridgoville	130	0.2	0.4	0.7	1.3	2.2	4.2	7.8	13	22
	Intervening - Ridgeville to Eaton	174	2.1	2.3	2.8	4.0	5.8 .033	10.8	17.2	25	36
16	Mississinewa River near Eaton	304	2.3	2.7	3.5	5.3	8.0	15	25	38	58
	Intervening - Eaton to Marion	373	20.7	23.3	26.5	31.7	37	47	60	82	.314
31	Mississincwa River at Marion (adjusted for low flow regulation)	677	23	26	30	37	45	62	85	120	175
47	White River at Muncie (adjusted to natural flow by adding diversion for water supply)	242	7.6	9.0	11,045	15	20	29	42	59	80
50	Buck Creek near Muncie	36.7	4.5	5.3	69.	8.0	10	13	16.436	19	22
	Intervening - Muncie to Anderson excluding Buck Creek	122	18	20	211	25	30	38	47	57	73 ,598
09	White Rivor at Anderson (adjusted to natural flow by correlation with White River near Noblesville)	401	30	34	38	48	60	80	105	135	175
	Intervening - Anderson to Noblesville	£13	29	31,075	35	42,102	50	60	80	110	155
78	White River near Noblesville	814	59	65	73	90	011	140	185	245	330

Table 5. -- Selected low-flow frequency and flow-duration data for partial-record stations and miscellaneous sites in the upper Mississinewa and White River basins embracing Delaware County (Data adjusted to period 1924-63 on basis of correlation with stream-flow records at stream-gaging stations).

Reference	Stream and location	Drainage	Annual low-flow, in cfs, for indicated period of con-secutive days and for indicated recurrence interval in years	in cfs, for od of con- und for in- unce interval,	Flow, in cfs, which was equaled or exceeded for indicated percent of time	ch was equaled indicated per-
•		(sq mt)	7 day		Ue e	(C)
			2 yr	10 yr	3	,
10	Campbells Creek near Desoto	20.4	0.2	0	0.3	0.2
19	Rees ditch near Eaton	13,5	1,1	ω	1.0	œ
22	Big Lick Creek near Wheeling	76.4	3.2	1.4	4.0	3.0
23	Studebaker ditch near Wheeling	10.5	8.	0	·:	4.
32	$^{-1}/$ White River near Harrisville	21.3	ι	8.	۳.	ю _.
37	Stoney Creek near Blountsville	18.7	1,4	ъ <u>.</u>	1.4	<u>.</u>
41	Cunningham ditch near Now Burlington	1,95	ρ	4.	φ.	w.
52	Bell Creck near Cross Roads	16,0	1.2	9·	1.2	αo
53	William Creek near Cross Roads	11.7	9.		œ.	ω,
54	No Name Creek near Progress	7,55	.2	0	2.	1,
57	York Prairie Creek near Cammack	9.70	2.6	1.8	2,4	2.0
63	Killbuck Creek near Bethel	26,1	2,0	oc.	1.8	1.2
29	Jakes Creek near Bethel	18.5	_.	0	ů,	и
70	1/ Killbuck Creek near Anderson	98.0	1.4	6.7	13	ග ් හ
91		44.7	9,6	1.6	3,4	2.4
		-				

-13-

Partical-record station 17

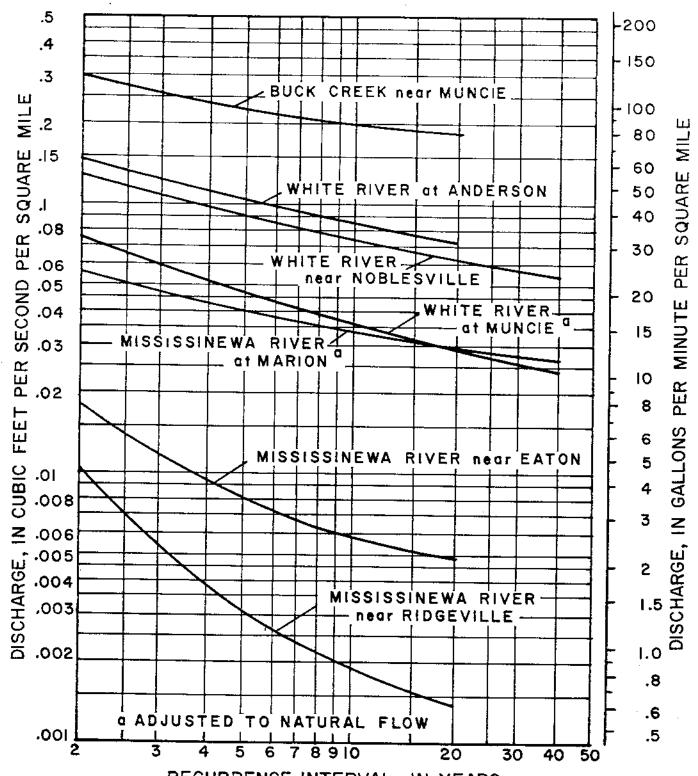
Adjusted to natural flow

Table 6 .-- Low-flow frequency as taken from stream-gaging station low-flow frequency curves (adjusted to the period 1924-63)

-					•																		•						····				Rei	
							47								31					····]6			 -					ယ		Reference	
					ing diversions for water supply)	(adjusted to natural flow by add-	White River at Muncie			-				(adjusted for low-flow regulation)	Mississinewa River at Marion								Mississinewa River near Eaton								Mississinewa River near Ridgeville		Station	
		·					242								677								304								130		area (sq mi)	Drainage
274	183	120	60	30	14	7	ω	274	183	120	60	30	14	7	ω	274	183	120	60	30	14	7	ω	274	183	120	60	30	14	7	သ		Consecutive days	
110	59	40	27	22	20	18	17	300	135	96	19	47	41	38	36	120	47	23	11	7.9	6.4	5.5	4.6	50	18	9.0	4.1	2.3	1.6	1.4	1.1	cfs	2	
.454	.244	.165	.112	.091	.083	.074	.070	.443	.199	.142	.090	.069	.060	.056	.053	.395	,155	.076	.036	.026	.021	.018	.015	,385	.138	.069	.032	.018	.012	.011	0.003	cfsm		
٥/	జ	24	17	14;	12	11	11	140	70	2	39	33	29	27	25	47	16	8.7	4.5	3.4	2.8	2.5	2.1	18	5.9	2.8	1.3	• &		.4	0.3	cfs	5	
. 236	.136	.099	.070	.058	.050	.045	.045	. 207	.103	.080	.058	.049	.043	.040	.037	. 155	.053	.029	.015	.011	.009	.008	.007	.138	.045	.022	.010	.006	.004	.003	0.002	cfsm		
42	25	19	13	11	9.3	8	8.3	96	54	42	32	28	24	22	21	28	9.7	5.7	3.0	2.5	2.0	1.8	1.5	10	3.3	1.6	.7		.	i.	0.2	cfs		Recurrence
.174	.103	.078	.054	.045	.038	.036	.034	.142	.080	.062	.047	.041	.035	.032	.031	.092	.032	.019	.010	.008	.007	.006	.005	.077	.025	.012	.005	.004	.002	.002	0.002	cfsm	10	1
1 34	20	15	10	8.6	7.4	7.0	6.7	74	44	35	26	24	21	20	19	18	6.6	4.1	2.3	2.0	1.7	1.5	1.3	თ	2.0	1.0		.4	. 2	. 22	0.2	cfs		Interval in
1 .140	.083	.062	.041	.036	.031	.029	.028	.109	.065	.052	.038	.035	.031	.030	.028	.059	.022	.013	.008	.007	.006	.005	.004	.048	.015	.008	.004	.003	.002	.002	0.002	cfsm	20	Years
31	18	14	9.3	7.6	6.5	6.2	5.9	64	39	32	24	22	20	19	18																	cfs	30	
.128	.074	.058	.038	.031	.027	.026	.024	.095	.058	.047	.035	.032	.030	.028	0.026				_													cism		
1 62	17	13	8.6	7.0	6.0	5.7	5.4	59	37	31	23	21	19	18	17																	cfs	40	
1 021.	.070	.054	.036	.029	.025	.024	.022	.087	.055	.046	.034	.031	.028	.026	0.025	. <u>.</u>																cfsm		

Table 6.--Low-flow frequency as taken from stream-gaging station low-flow frequency curves (adjusted to the period 1924-63)

												 -			<u> </u>					 -						
							78								60								50	no.	Reference	
							White River near Noblesville					near Noblesville)	correlation with White River		at								Buck Creek near Muncie	Station		
							814								401								36.7	area (sq mi)	Drainage	
2/4	183	120		6 J	20 14	7	<u>ယ</u>	274	183	120	60	30	14	7	ယ	274	183	120	60	30	14	7	ယ	Days	Consecutive	
410	240	100	140	125	110	105	99	220	130	100	77	67	61	58	5 4	24	18	16	13	12	11	11	10	cfs	2	
.504				. 154	. 135	.129	.122	.549	.324	.249	.192	.167	.152	.145	.135	.654	.490	.436	.354	.327	.300	.300	0.272	cfsm		
230	145	4115	97	85	76	73	68	125	80	68	54	48	43	41	38	16	14	12	10	9.6	8.8	8.4	8.1	cfs		
. 282	.178	.141	.119	.104	.093	.090	.084	.312	.200	.170	.135	.120	.107	.102	.095	.436	.382	.327	.272	.262	.240	.229	0.221	cfsm	יי	
165	115	95	79	70	62	60	56	91	64	56	44	39	36	34	32	13	12	10	9.0	8.4	7.8	7.4	7.0	cfs		Recurrence
. 203	.141	,117	.097	.086	.076	.074	.069	. 227	.160	.140	.110	.097	.090	.085	.080	.354	.327	.272	.245	.229	.212	. 202	0.191	cfsm		- 1
130	99	82	66	59	53	51	48	72	54	48	38	33	30	29	28	12	11	9.5	8.2	7.6	7.0	6.7	6.4	cfs 20	Т,	Interval in
.160	.122	.101	.081	.072	.065	.063	.059	.180	.135	.120	.095	.082	.075	.072	.070	.327	.300	. 259	. 223	.207	.191	.183	0.174	cfsm		Years
120	91	75	60	53	49	47	44																	cfs 30		
.147	.112	.092	.074	.065	.061	.058	0.054								·									0 cfsm		
110	86	71	57	50	46	44	4 22										_			: 1 %				cfs		
.135	.106	.087	.070	.061	.056	.054	0.052																,	40 cfsm		



RECURRENCE INTERVAL, IN YEARS
Figure 3-Magnitude and frequency of 7-day low
flow at seven gaging stations on the Mississinewa
River, White River, and Buck Creek
(adjusted to period, 1924-62)

Base-flow measurements. The third technique used to analyze the quantity of water available in the streams and also to define areas of good or poor ground water possibilities is the base-flow measurement run. In this technique numerous discharge measurements are made over a short period of time throughout an area so that comparisons of flow in streams may be made at approximately equal flow durations. This method is useful in assessing the relative yields of different drainage basins. A word of caution is needed. Because there is only one measurement of flow made at one particular time, there is the possibility that the discharge may be affected by some temporary or permanent work of man. This may be in the form of diversions, depletions, or regulations of the streamflow, or addition to the flow by sewage effluents. Such alterations are not always readily apparent in a reconnaissance type survey which involves collection of base-flow measurements. Also, streams in a certain area may be severely affected by a series of extreme hydrologic events.

A group of 73 base-flow measurements were made the week of November 11, 1963, when the base flow throughout the study area was very close to the 90 percent flow duration. Although most measurements were made within Delaware County, they were not limited to this political unit. Measurements were made in the entire drainage system embracing the county. The measured discharges have been shown on plate 1, as well as the flow in cfsm of areas between the measured sites.

Draft-storage requirements. The natural flow of a stream during low flow is often not sufficient to meet all the demands of water supply. It then becomes necessary to search out other ways of overcoming the deficiency. One solution to the problem is storing water during periods of excess flow and releasing it during times of deficient flow. In order to determine the amount of storage necessary to maintain certain draft rates, storage-required graphs for various frequency recurrence intervals have been computed from the low-flow frequency data and shown as figures 4 and 5 for the following gaging stations: Mississinewa River near Ridgeville, Mississinewa River near Eaton, White River at Muncie, White River at Anderson and Buck Creek near Muncie.

In using the storage-required frequency curves for estimates of storage, it must be kept in mind that such graphs have not been corrected for reservoir seepage and evaporation.

Summary of Low-flow Stream Characteristics

Results obtained by use of flow duration and low-flow frequency analyses and base-flow discharge measurements show that the low-flow yield per square mile of the Mississinewa and White Rivers generally increases progressively downstream as they pass through Delaware County. Also, low flow is better sustained in the White River than in the Mississinewa River. The low flows of tributary streams are variable at about 90 percent flow duration and range from no flow for most streams in the northern half of the county having less than 10 square miles of drainage area, including the headwaters of larger streams, to the

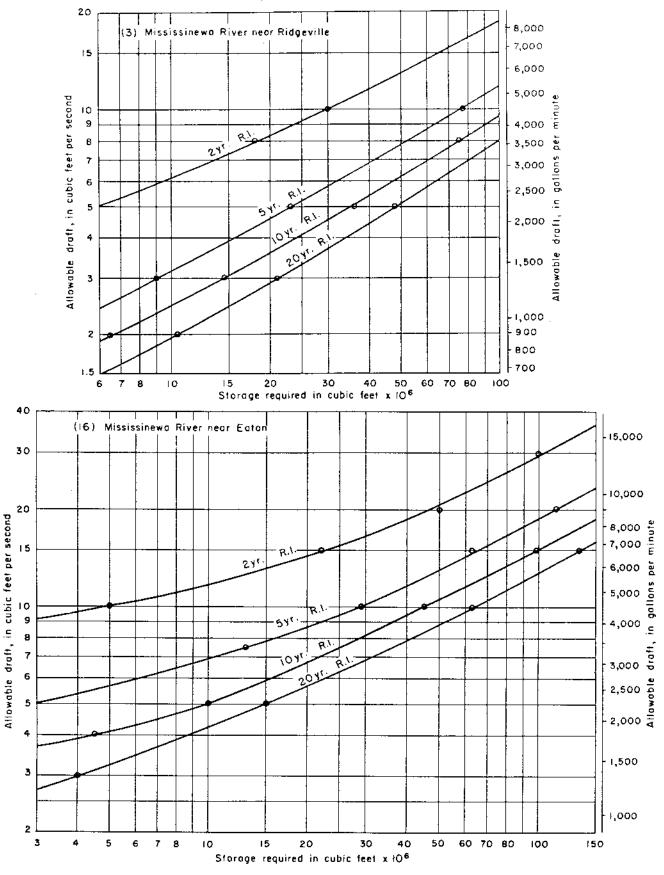


Figure 4.-- Storage-required graphs for 2, 5, 10 and 20-year recurrence intervals for the Mississinewa River near Ridgeville and Mississinewa River near Eaton. (Storage is uncorrected for reservoir seepage and evaporation)

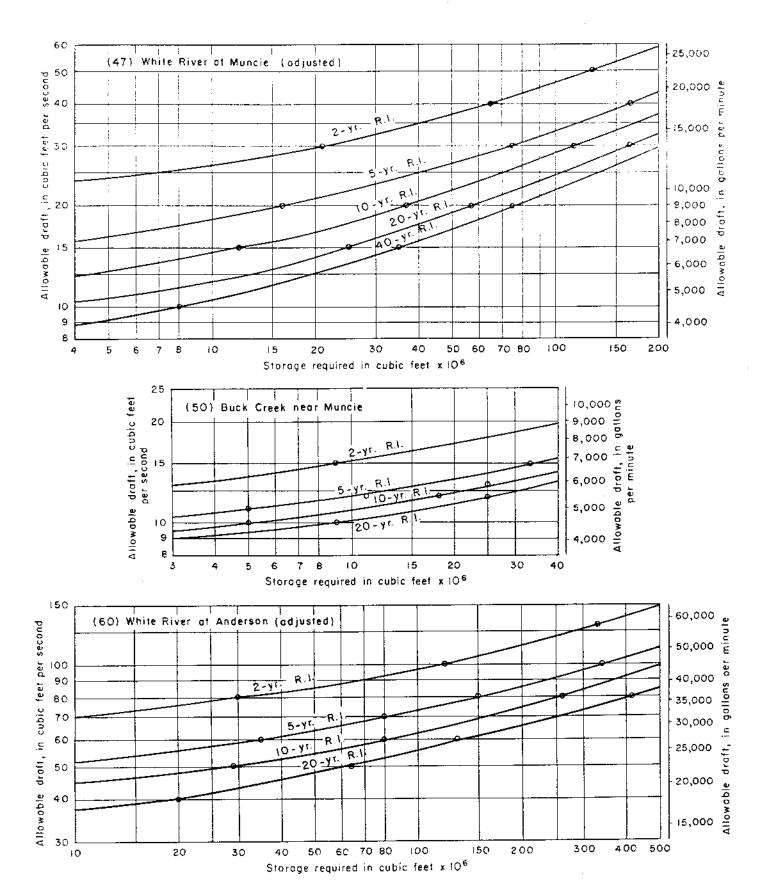


Figure 5 .-- Storage-required graphs for 2,5,10,20 and 40-year recourence intervals for the White River at Muncie and 2,5,10 and 20-year recurrence intervals for Buck Creek near Muncie and White River at Anderson . (Storage is uncorrected for reservoir seepage and evaporation.

excellent sustained flow of 0.26 cfsm (0.17 mgd per square mile) for the upper part of Buck Creek (Plate 1).

The flow-duration and low-flow frequency data for the 7 stream-gaging stations show Buck Creek near Muncie to have the highest sustained low-flow yield per square mile and the Mississinewa River near Ridgeville to have the lowest. The relatively flat slopes of the duration and frequency curves, with the exception of that for the Ridgeville station, indicate that a considerable quantity of ground-water discharge is available to sustain low flow.

At the 90 percent flow duration, the flow in the Mississinewa River at Marion is 2.5 times as great per square mile as that upstream near Eaton, and 3.9 times as great as that further upstream near Ridgeville. The spread is even larger at the 95 and 98 percent duration values. Data in table 4 indicate that the flow from the intervening area between the gages at Marion and near Eaton at 90 percent flow duration is 3.0 times as great as the intervening area between the gages near Eaton and near Ridgeville and 5.8 times as great as the area upstream from the near Ridgeville gage.

Carrying the comparison to the White River at 90 percent flow duration, the natural (adjusted) flow per square mile of the White River at Muncie is 3.2 times as great as the Mississinewa River near Eaton, but 0.8 as large as the flow from the intervening area between the Eaton and Marion gages. The 90 percent flow duration from the intervening area between the White River at Anderson and the White River at Muncie, excluding the area above the Buck Creek near Muncie gage, is 3.0 times the flow contributed from the area above the Muncie gage, but 0.9 times the flow from the area above the Buck Creek near Muncie gage. Although the drainage area for the stream-gaging station on Buck Creek near Muncie equals only 15 percent of the drainage area above the stream-gaging station on White River at Muncie, the discharge is 50 percent as large at 90 percent flow duration.

The low-flow characteristics at the 7 stream-gaging stations may be examined further by analyzing the minimim observed flows that have occurred during their period of record. Table 7 shows the instantaneous minimum flow, and the lowest two minimum average flows for 1-day, 3-day, 7-day, 14-day, 30-day and 60-day consecutive periods with the beginning month of the period. The values in this table when compared with the low-flow frequency data in table 6 make it possible to estimate the severity of the observed low flows during the period of records, excepting for the stream-gaging stations that have pronounced regulation or diversion. Example: The lowest 3-day average flow on the Mississinewa River near Eaton during the period from 1952-63 is 2.1 cfs (table 7). This flow will occur on the average 1 time in every 5 years (table 6).

In reviewing the low-flow characteristics it is interesting to compare the runoff map (Plate 1) with the shape of the piezometric water surfaces in the glacial drift (fig. 8) and bedrock (fig. 13). Where contours on the piezometric surface are parallel to the streams, flow is sustained by ground-water discharge. It is possible that some of the

flow does not appear in the streams but rather occurs as underflow beneath the stream channels. This may be the case in the reach of the White River from Yorktown to Daleville.

Application of Techniques to Water Planning

Application of the four foregoing techniques to water planning is demonstrated by the following hypothetical example. A new industrial plant desires to locate on the Mississinewa River at Eaton, where the drainage area is about 310 square miles. It is determined that the water supply obtainable from the town of Eaton's wells is more than adequate to meet anticipated future needs. However, in the absence of a sewage disposal plant at Eaton, it will be necessary to have a minimum flow of 6 cfs, or 3.9 mgd, for dilution of plant waste products as predetermined by the Indiana Stream Pollution Control Board.

Observed minimum flows (at the gage upstream near Eaton) (table 7) are less than the required flows; therefore, it is necessary to know the probable percent of time there will be a shortage of water. From figure 2, it is determined that the flow of 6 cfs, or 0.02 cfsm, will be available about 93 percent of the time, or about 7 percent of the time the flow would be less than that required for dilution.

A management decision could be made at this point as to whether or not plant operation can be economically reduced or shut down 7 percent of the time.

If management decides that plant operation cannot be shut down 7 percent of the time, then arrangements for storage or a supplemental supply must be considered and planned.

To assist further in the water planning, the low-flow frequency data for the gaging station near Eaton indicates the frequency at which the discharge falls below 6 cfs. Table 6 shows that on the average, about once every 2 years for a 14-day consecutive period, or once every 10 years for a 120-day consecutive period, or once every 20 years for a 183-day consecutive period, the average natural flow in the Mississinewa River would be less than that needed for plant operations. To continue the plant in operation during these periods of insufficient streamflow, it would be necessary to supplement the natural streamflow with dilution water from a standby facility such as private wells, the Eaton water supply system, or upstream surface storage on the Mississinewa River or on one of its tributaries.

The base flow measurements would be used to determine which of the Mississinewa River tributaries might offer the best potential for off channel storage should it be decided to go to a small storage dam to supplement the natural flow of the Mississinewa River. An examination of plate 1 indicates that both Bosman and Rees ditches offer good possibilities for low-flow yield in their lower reaches.

Table 7.--Average discharge, instantaneous minimum, and lowest two minimum average flows for 1-, 3-, 7-, 14-, 30-, and 60-day consecutive periods for stream-gaging stations on the Mississinewa River, White River, and Buck Creek

	· · ·								<u>-</u>						T
	78			60		50		47		31		16		ယ	Reference
	White River near Noblesville			White River at Anderson		Buck Creek near Muncie		White River at Muncle		Wississinewa River at Marion		Mississinewa River near Eaton		Mississinewa River near Ridgeville	Station
	814			401		36.7		242		677		304		130	Drainage area (sq mi)
Sept. 1926, Oct. 1928 to Sept. 1963	May 1915 to	Sept. 1963	Sept. 1926, Oct. 1931 to	July 1925 to	Sept. 1963	Oct. 1954 to	Sept. 1963	Nov. 1930 to	Sept. 1963	Sept. 1923 to	Sept. 1963	Mar. 1952 to	Sept. 1963	Aug. 1946 to	Period of record
	803			375		35.9		b 217		641		282	•	132	Average discharge (cfs)
	36	***************************************		8.8		-		.6		م د	-11	2.0	·,	0.1	
	(9/25/41)			(9/24/40)				(9/16/37)		(4/17/59)		(9/23,27/54)		(10/24/46)	Instantaneous minimum (cfs)
46 (9/40)	39 (9/41)		12.0 (9/41)	° 9.1 (9/40)	6.0 (12/62)	5,8 (9/56)	1.1 (10/56)	1.1 (9/54)	3.8 (10/43)	3.8 (10/40)	2.2 (10/53)	2.0 (9/54)	.3 (9/54)	0.1 (10/46)	1-day (cfs)
51.0 (9/40)	41.3 (9/41)		13.0 (9/40)	c 12.3 (9/41)	6.3 (12/62)	6.1 (9/56)	1,3 d	1.1 (9/54)	5.4 (10/40)	2 5.1 (10/43)	2.2 (10/53)	2.1 (9/54)	.3 (9/54)	0.2 (10/46)	3-day (cfs)
53.1 (9/40)			15.7 (9/40)	13.0 (9/41)	7.7 (9/60)	e 7.3 (12/62)	1.4 (8.34)	c 1.2 (9/54)	12.0 (11/28)	8.4 (10/40)	2.5 (9/53)	2.2 (9/54)	.3 (9/54)	0.2 (10/46)	7-day (cfs)
53.1 (9/40)	43.9 (9/41)		18.9 (9/40)	c 13.9 (9/41)	8,6 (9/56)	8.0 (9/60)	1.4 (7.34)	c 1.2 (9/54)	18.9 (10/40)	a 12.0 (11/28)	2.5 (9/53)	2.3 (9/54)	.5 (9/54)	0.3 (10/46)	14-day (cfs)
57.2 (55.1 (20.8 (20.5	9.4 (9.1 (2.2	1.5	21.7	a 16.4 ()	2.9	2.3	1.0	0,6	30-day
(9/40) 6:	(9/41) 5		(9/41) 2	(9/40) 2	(8/60) 1	(9/56)	(9/56) c	(9/54) c	(8/40) 2	(10/28) a	(9/53)	(9/54)	(9/46)	(9/54)	4
62.6 (8/40)	59.8 (8/			22.1 (8/	10.3 (8/	9.8 (9/	&	2.9	22.8 (8/	20.2 (9)	6.8 (8,	4.3 (8,	1.7 (9,	0.9 (9,	60-day (cfs)
40)	(8/41)		(8/40)	(8/41)	(8/60)	(9/56)	(8/40)	(9/56)	(8/40)	(9/28)	(8/54)	(8/53)	(9/53)	(9/46)	

Regulated

ъ,

Adjusted for diversion after September 1937

Diversion for water supply above gage, returned to stream through sewage treatment plant below gage

^{8/34} and 10/56

Ice effect

To determine the approximate amount of storage necessary to maintain a 6 cfs draft rate from a channel-storage reservoir on the Mississinewa River itself, storage-required frequency curves for the near Eaton gage as presented in figure 4 may be used. A reservoir of 16 million cubic feet, or 120 million gallons, capacity would be needed to maintain a 6 cfs flow for a 10-year frequency recurrence interval. To maintain this draft rate to overcome a 20-year frequency drought, 23 million cubic feet, or 172 million gallons, of storage would be necessary. Additional storage would be necessary to overcome water loss by seepage and evaporation.

Development Potential

The available water supply potentials of the several streams in Delaware County are an index of their development potential. Many of the streams flow with fairly well sustained low flows during dry weather. An excellent example of development is Prairie Creek, a tributary to White River, which was put to use in 1961 for water supply by the Muncie Water Works Company by construction of a 7.2 billion gallon reservoir.

The variation in low flow is illustrated by plate 1. The color lines along the streams indicate the range of flow into which the measured discharges fell during the week of November 11, 1963, at about 90 percent flow duration. The color lines between the measured locations are estimates based on flows at measured points.

From a practical standpoint, the low flow available 90 percent of the time may be considered the maximum dependable supply that can be obtained without storage in reservoirs. The other 10 percent of the time the flow is less than this amount.

The average discharges for the period of record at each of the 7 stream-gaging stations are included in table 7. The average discharge divided by the drainage area equals about 1 cfsm at each gage. The average discharge represents the absolute limit of the potential surfacewater supply available, if it were possible to store in reservoirs and release at uniform rate the total water that flows from the area above the gage. To carry this reasoning one step further to ungaged sites, an estimate of the absolute limit of the potential surface-water supply available on any stream in Delaware County, regardless of size, would be the product of the drainage area, in square miles, times 1 cfsm.

Mississinewa River and tributaries. The flow in the Mississinewa River increases from a yield of 0.008 cfsm at the Ohio-Indiana state line, to 0.03 cfsm, at about 90 percent flow duration before it leaves Delaware County (pl. 1). On the basis of the streamflow data, the Mississinewa River offers a better development potential downstream from about the mouth of Campbells Creek than it does above this point. This increase in base flow is due to better yields of ground-water discharge from (1) the till and (2) probably the valley train and outwash plain sediments beginning on the main channel a little upstream from Eaton and continuing downstream. Another feature of the surficial

geology in the area of increased base flow yields is the presence of large sand and gravel deposits in the forms of kames and eskers in the till.

A large user of water considering the Mississinewa River as a supply will have to depend on an impoundment if his activity requires 9 mgd or more, more than 90 percent of the time.

The tributaries in Delaware County offer various development potentials at about 90 percent flow duration. Mud Creek, Halfway Creek, the upper reaches of Bosman ditch, Rees ditch, Pike Creek, as well as a few smaller unnamed tributaries, were either dry or contained very little flow during the week of November 11, 1963. Some of these streams had excellent amounts of flow for their size in their lower reaches.

Most of the tributaries having less than 10 square miles of drainage area, including the headwaters of larger streams, go dry frequently. Impoundments would be necessary if these sources were developed. If a supply of 0.6 mgd, or more, was required more than 90 percent of the time from any of the aforementioned tributaries, an impoundment would be necessary.

White River and tributaries. The flow in the White River increases from a yield of 0.03 cfsm in its headwaters in Randolph County to 0.10 cfsm in Delaware County at about 90 percent flow duration (pl. 1). This increase in base flow is due to better yields of ground-water discharge from (1) the till and (2) probably the valley train and outwash plain sediments, beginning on the main channel of White River at about Mud Creek and continuing downstream.

Below Muncie, the White River meanders through the outwash sediments, and although the flow in the reach from below Muncie to the county line nearly doubles at 90 percent flow duration, the increase is due largely to the entrance of Buck Creek and York Prairie Creek. The White River from below Mud Creek to the mouth of Buck Creek would offer the best potential development under natural flow conditions. However, the flow in this reach is altered by the water supply for the city of Muncie. Water is taken out above Muncie and sewage effluent is returned below Muncie.

A large user of water considering the White River as a supply will have to depend on an impoundment if his activity requires 25 mgd or more, more than 90 percent of the time.

The tributaries in Delaware County offer various development potentials at about 90 percent flow duration. The upper portions of Muncie Creek, and some of the small tributaries of Killbuck and Pipe Creek were either dry or contained practically no flow during the week of November 11, 1963. Flows in Stoney Creek, Mud Creek, Medford drain, the middle portion of Muncie Creek, Bell Creek, the upper portions of Killbuck Creek and Pipe Creek are fairly well sustained. The tributaries of the White River offering the best potential development are Cunningham ditch (tributary to Prairie Creek Reservoir and thus already used in water

supply development), Buck Creek (used for water supply by the Muncie Water Works Company without impoundment), York Prairie Creek and Brandon ditch (tributary to Fall Creek).

Most of the tributaries in the northern half of the county having less than 10 square miles of drainage area, including the headwaters of larger streams, go dry frequently. Impoundments would be necessary if these sources were developed. Most of the tributaries in the southern half of the county, regardless of size, can be depended on to yield more than 9 gpm (0.013 mgd) per square mile 90 percent of the time. Buck Creek, in the upper part of its basin, can be depended on to yield more than 120 gpm (0.170 mgd) per square mile 90 percent of the time.

Glacial Deposits

Delaware County is underlain by glacial drift that ranges in thickness from 0 to about 300 feet. The principal sources of ground water in the drift are two sand and gravel aquifers (fig. 6). Nearly all the ground water in these deposits must percolate through the confining layer that covers almost the entire land surface of the county (fig. 6). Some water can be recovered by dug wells in the surficial confining layer, but usually these wells are not adequate even for domestic supplies.

Sand and Gravel Aquifers

One of the two important water-bearing zones underlies nearly the entire county and is referred to as the "principal Pleistocene aquifer" (fig. 6). The other zone lies beneath the principal Pleistocene aquifer, usually within or near buried bedrock valleys and is referred to as the "deeper sand and gravel aquifers".

The principal Pleistocene aquifer consists chiefly of sand and gravel as much as 80 feet thick, with some areas of sand alone. It extends beyond the boundaries of Delaware County, and is an important aquifer in surrounding areas. The aquifer is artesian except for a few scattered areas of water-table conditions in the central part of the county. Its stratigraphic position is shown on figure 6.

The deeper sand and gravel aquifers occur principally in the southern third of the county in the thicker parts of the glacial drift. The thickness and configuration of these aquifers are not well known.

Recharge and Discharge

Most of the inflow to the shallow glacial deposits of Delaware County is derived from precipitation within the county. Some water enters the county by underground flow, most of which occurs along the eastern and southern boundaries. Average inflow and outflow in the shallow glacial deposits are symbolized by figure 7. Recharge and discharge of the deeper sand and gravel aquifers are illustrated along with the Silurian Dolomite (p. 34).

		Symbol on cross section	. Lithology	Thickness (feet)	Areal extent	Importance for water supply
		Confining layer	Clay and silt; sandy, gravelly, calcareous, brown to yellowish-brown, occasionally blue. Mostly tilt	0-110	Underlies entire county except for a few very small isolated areas	A few dug wells for livestock. Limits re- charge to aquifers and ground-water dis- charge to streams
	undifferentialed	Aquifer	Sand and gravel; silty, often clayey, tan to brown	0-40	Underlies most of the county. Cannot be distinguished from principal Pleistocene aquifer where the till below is missing	Some domestic and stock wells
QUAYERNARY		Confining layer	Clay; silty, sandy, gravelly, usually brown or blue. Till	0-100	Same as above unit	Confining layer for part of principal Pleistocene aquifer
QUA	뎔	Aquifer	Sand and gravel; silty, often clayey	0-100	Underlies the entire county (fig. 9)	Principal Pleistocene aquifer. Domestic and industrial use
	Pleistocene a	Confining layer	Clay; silty, sandy, gravelly, usually blue or brown, sometimes red in lower part; with occasional sand and gravel lenses. Till	0 -240	Underlies most of the county. Missing on bedrock highs	Confining layer for most of the Silurian Dolomite
i		Aquifer	Sand and gravel; often interbedded with clay	0-70+	Occurs in and near most buried valleys (fig. 11)	Domestic and industrial use
	SILURIAN	Aquifer	Mostly determite. Drillers occasionally re- port shale and, rarely, sandstone	200-500	Underlies the entire county	Domestic, industrial, and municipal use

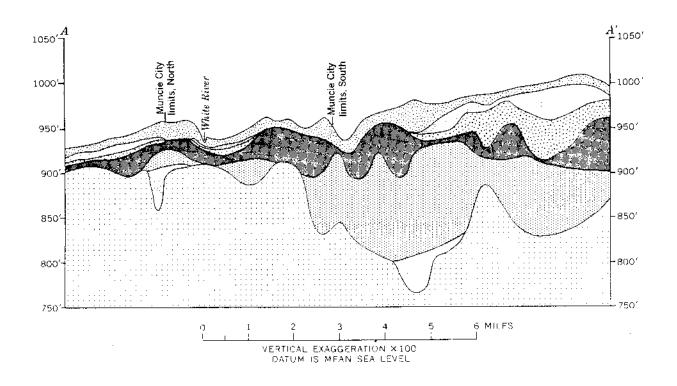
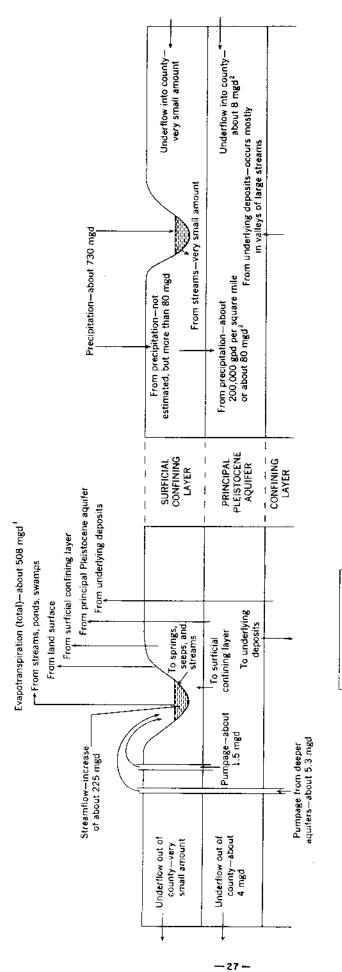


Figure 6.—Generalized cross section showing aquifers and confining layers of Delaware County. Line of cross section is shown on figure 9.



Increase of streamflow about 225 Evapotranspiration about 509¹ Underflow about 4 Total about 738		Average amount (mgd)
Evapotranspiration about 509 the Underflow about 4 Total about 738	Increase of streamflow	about 225
Underflow about 4 Total about 738	Evapotranspiration	about 5091
Total about 738	Underflow	about 4
	Total	about 738

Precipitation about 730
Underflow about 8²
Total about 738

Average amount (mgd)

2 See text

Figure 7.—Average recharge and discharge of the shallow glacial deposits.